

DEEP IMPACT

on a Super-Earth in the Vicinity of a Central Star

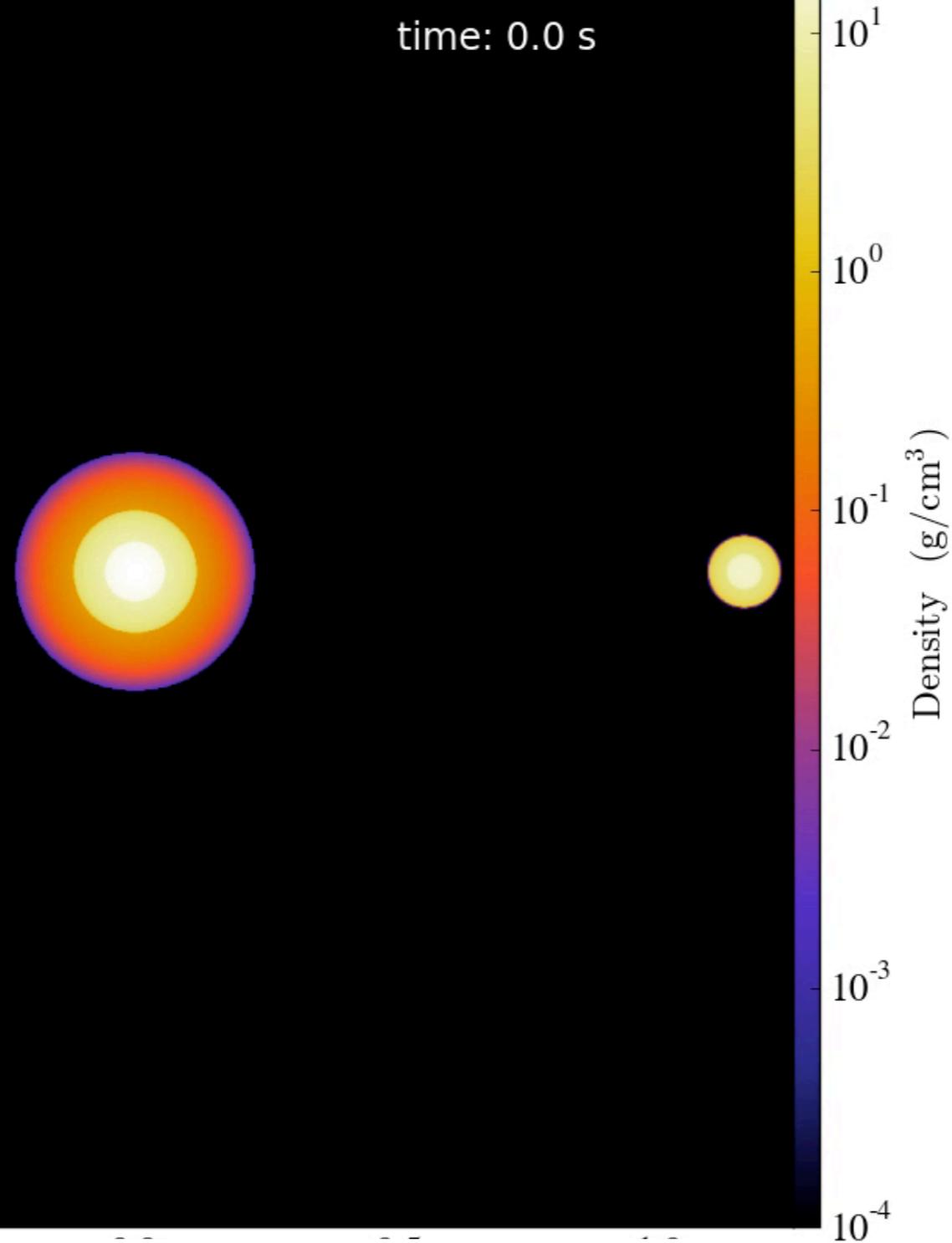
Yasunori Hori

*National Astronomical Observatory of Japan
Astrobiology center,
National Institutes of Natural Sciences*

Collaborators: **Shang-fei Liu (UCSC)
Douglas N.C. Lin (UCSC)
Erik Asphaug (Arizona State Univ.)**

What We did on Cray XC 30 : Giant Impact Simulations

(Liu, YH, Lin, & Asphaug, 2015)



XC-B user

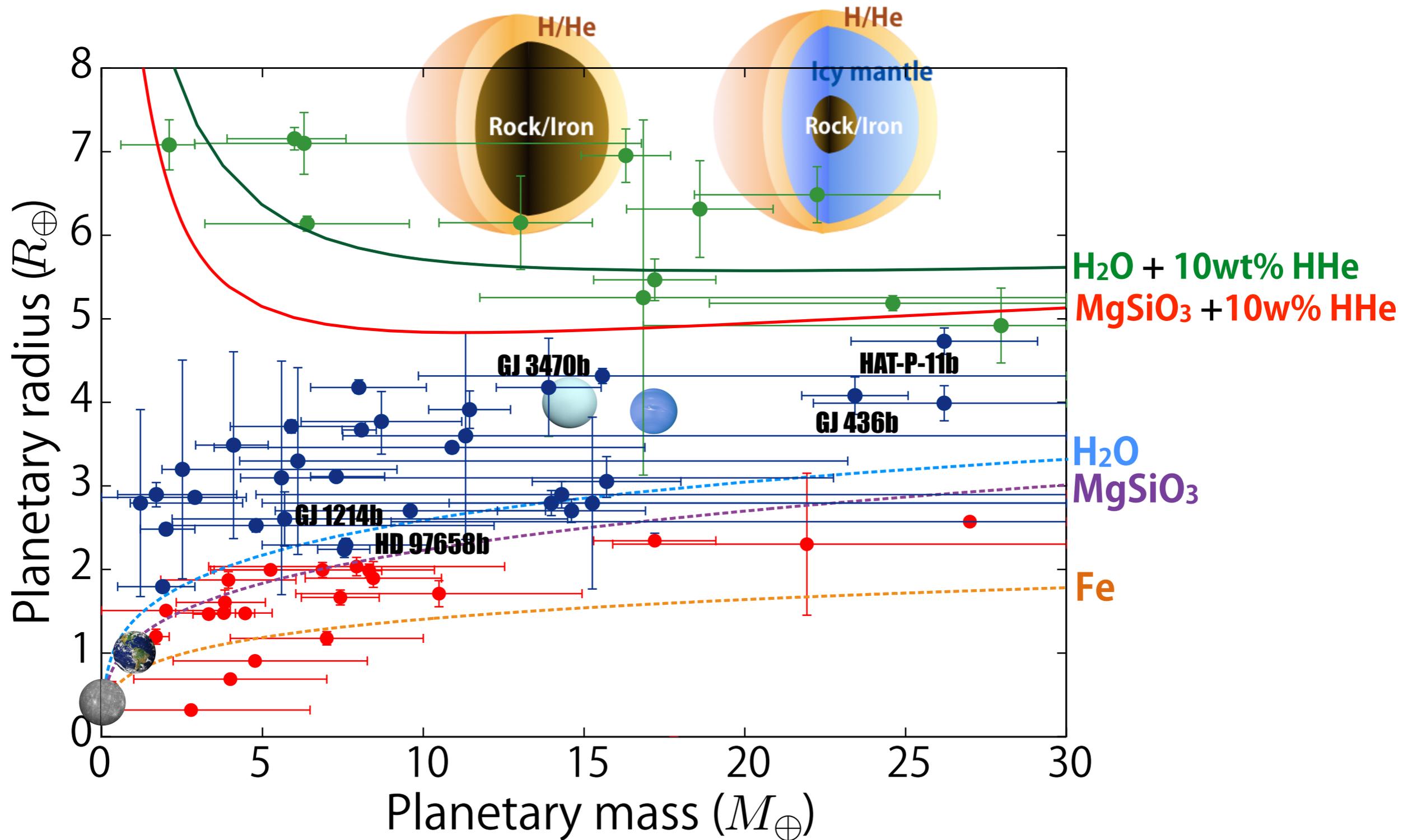
- 100-200 CPUs per run
- 2-3 weeks



1 paper **published**
2 papers **in prep.**

Compositional Diversity of Low-mass Planets

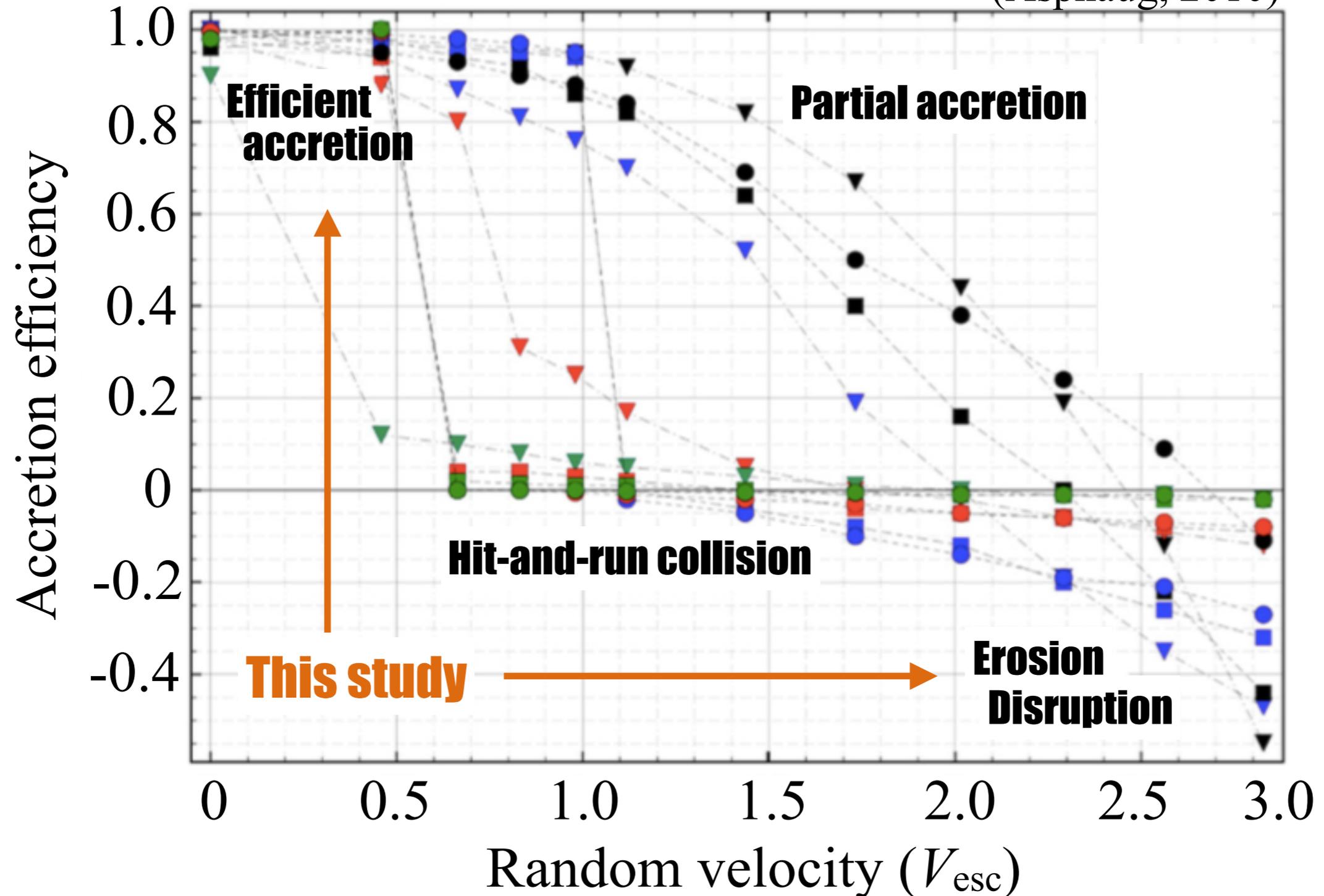
Mass-radius relationship of transiting planets with mass of $< 30 M_{\oplus}$



Giant Impacts: Accretionary and Destructive

Accretion efficiency as a function of mass ratio (0.1 ▼, 0.5 ■, 1.0 ●), impact angle (0, 30, 45, 60°), and impact velocity

(Asphaug, 2010)

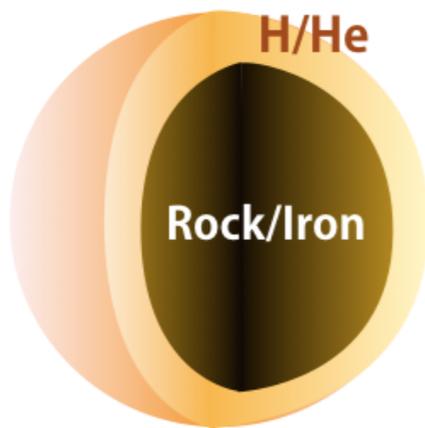


Modeling of a Giant Impact

Three-dimensional hydrodynamic simulations : FLASH with the AMR

- A pair of planets' center of mass frame (Fryxell *et al.*, 2000)
- The width of a computational domain ~ 1 AU
- include the tidal force from a central star
- impose an open boundary condition

Three-Layered interior structures of a target and an impactor



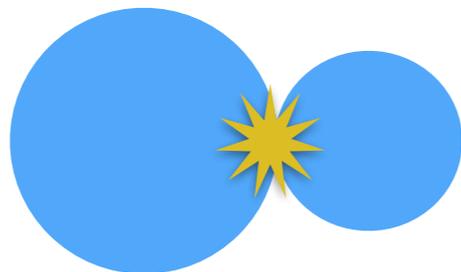
Tillotson EoS for rocky and iron material (Melosh, 1989)
rock (silicate) : iron = 2:1

Only a target has an atmosphere (7.5wt%)

Polytropic-type EoS for H/He gas ($H_2 : He = 7 : 3$)

(Liu *et al.*, 2013)

Giant impacts (@ 0.1 AU)



head-on collision

(1) Low-speed model (accretion regime) : $V_{\text{imp}} = V_{\text{esc}}$

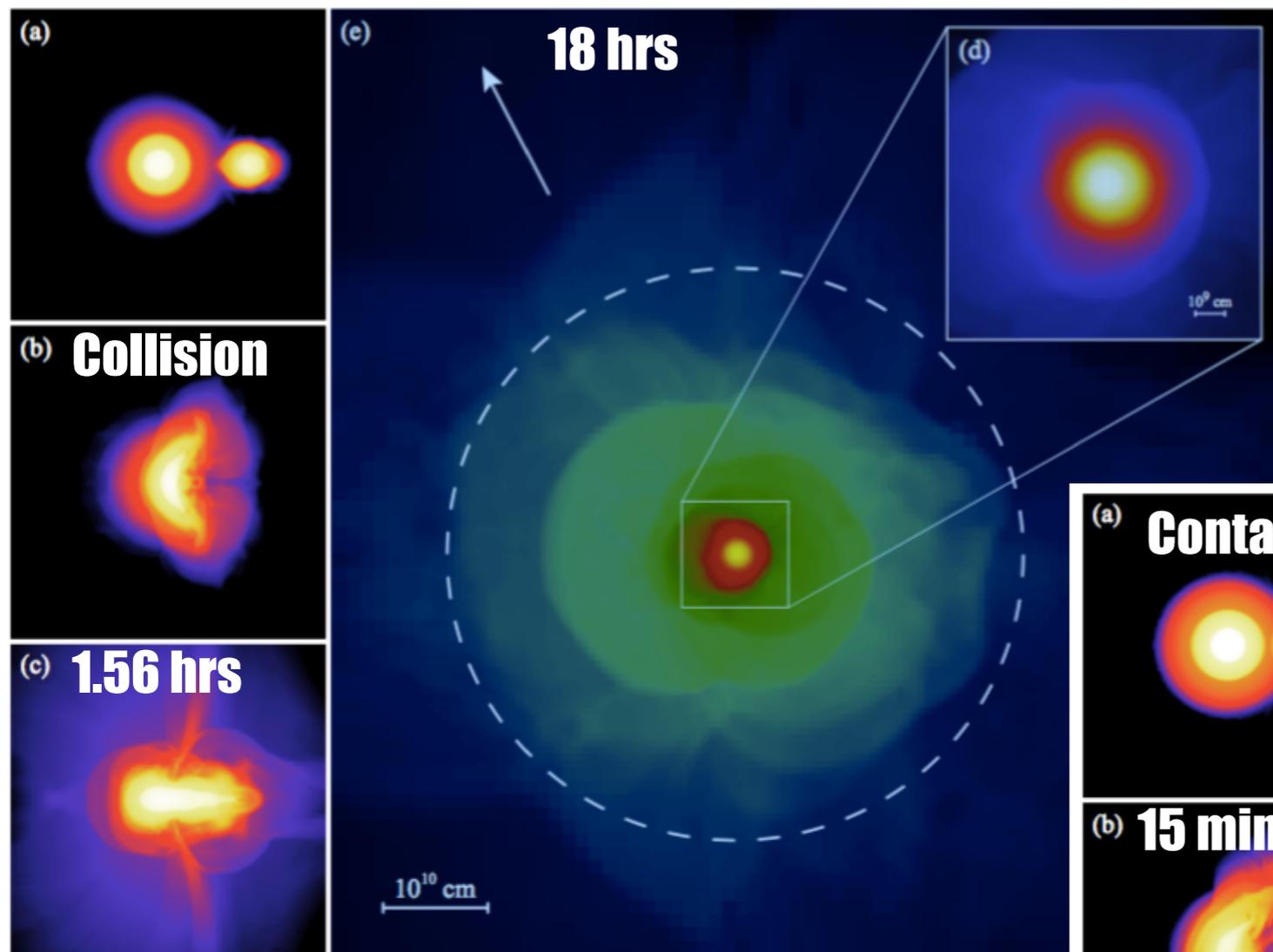
$4.3 M_{\oplus}$ & $1.0 M_{\oplus}$

(2) high-speed model (destructive regime) : $V_{\text{imp}} = 3V_{\text{esc}}$

$10 M_{\oplus}$ & $1.0 M_{\oplus}$

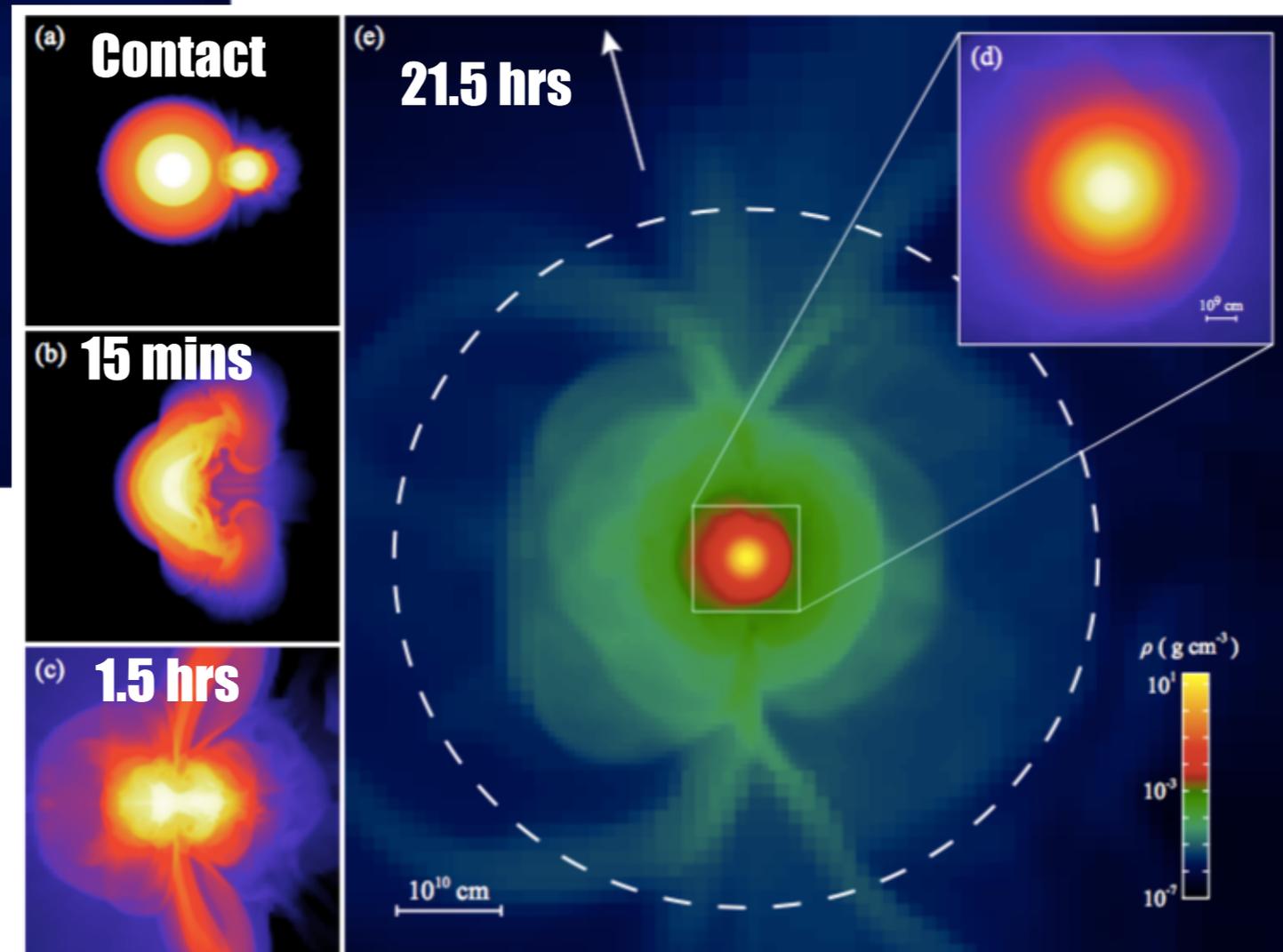
Snapshots of Two Head-On Collisions : Density Contours

(Liu, YH, Lin, & Asphaug, 2015)



Low-speed impact

~30% of the initial atmospheric mass evaporated via an impact



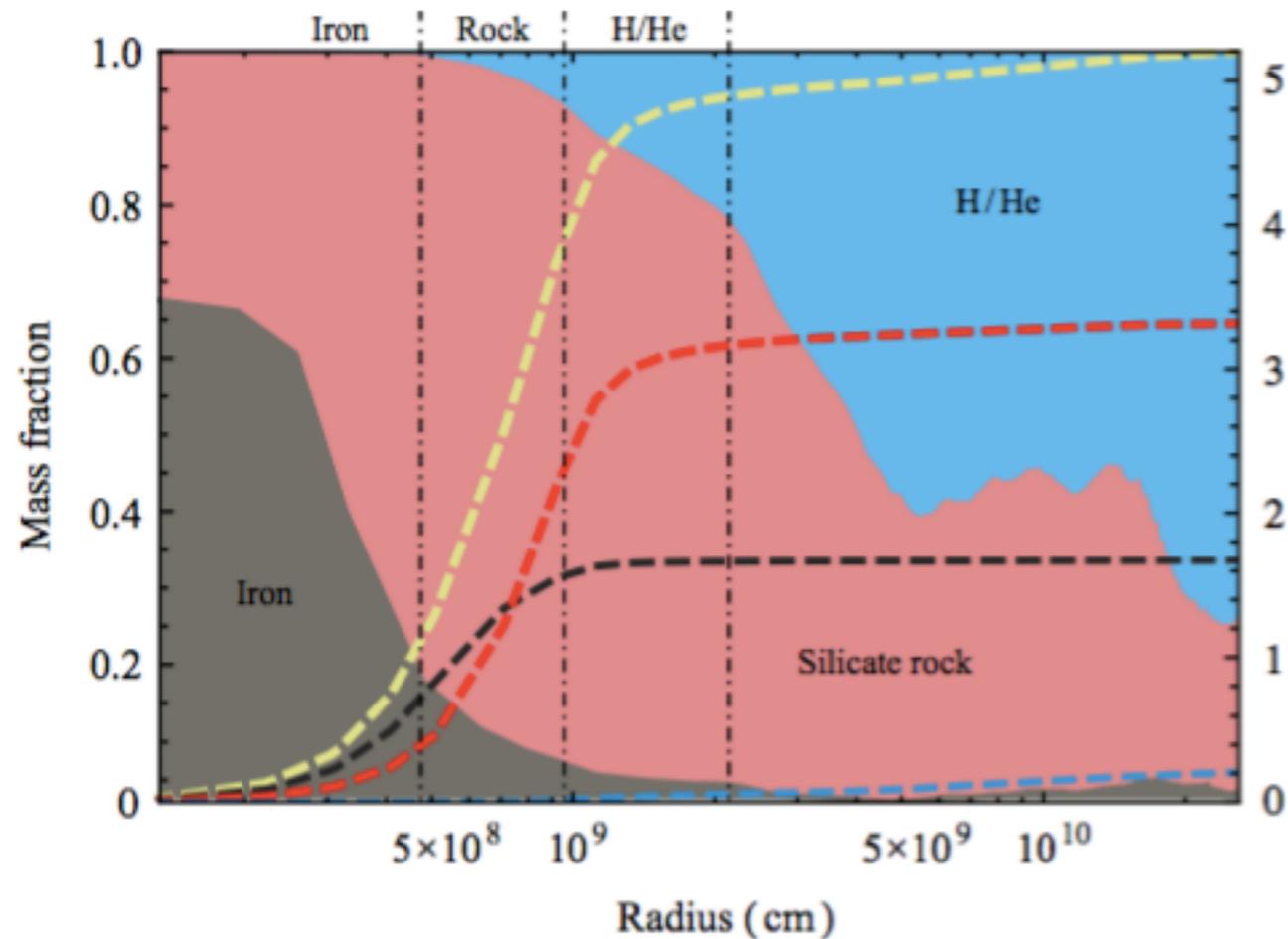
High-speed impact

The atmosphere is lost by ~80%

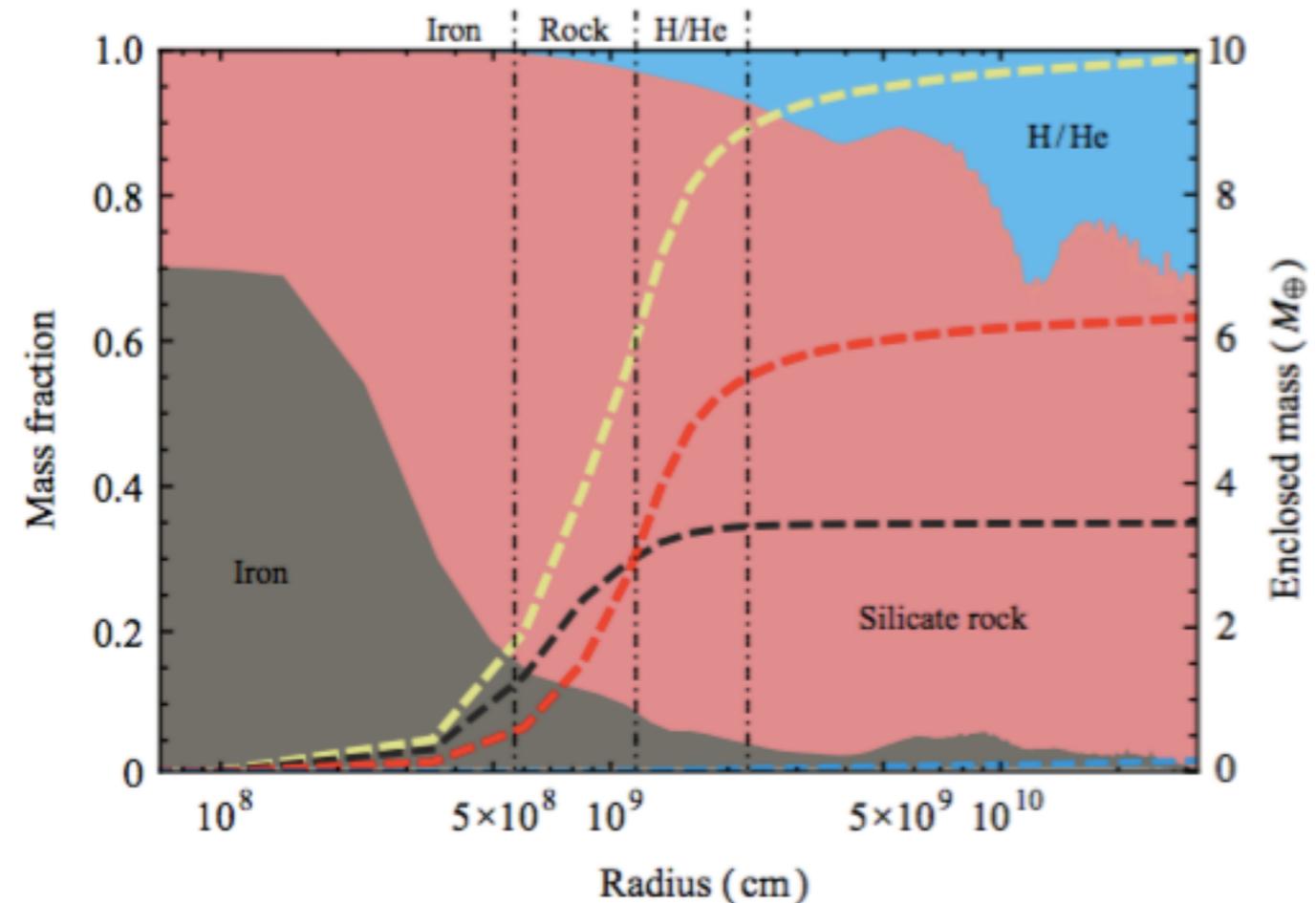
A hot atmosphere extends beyond the Hill radius and continues to lose via **the Roche-lobe overflow**

Radial Distribution of Each Species After a Collision

Low-speed impact

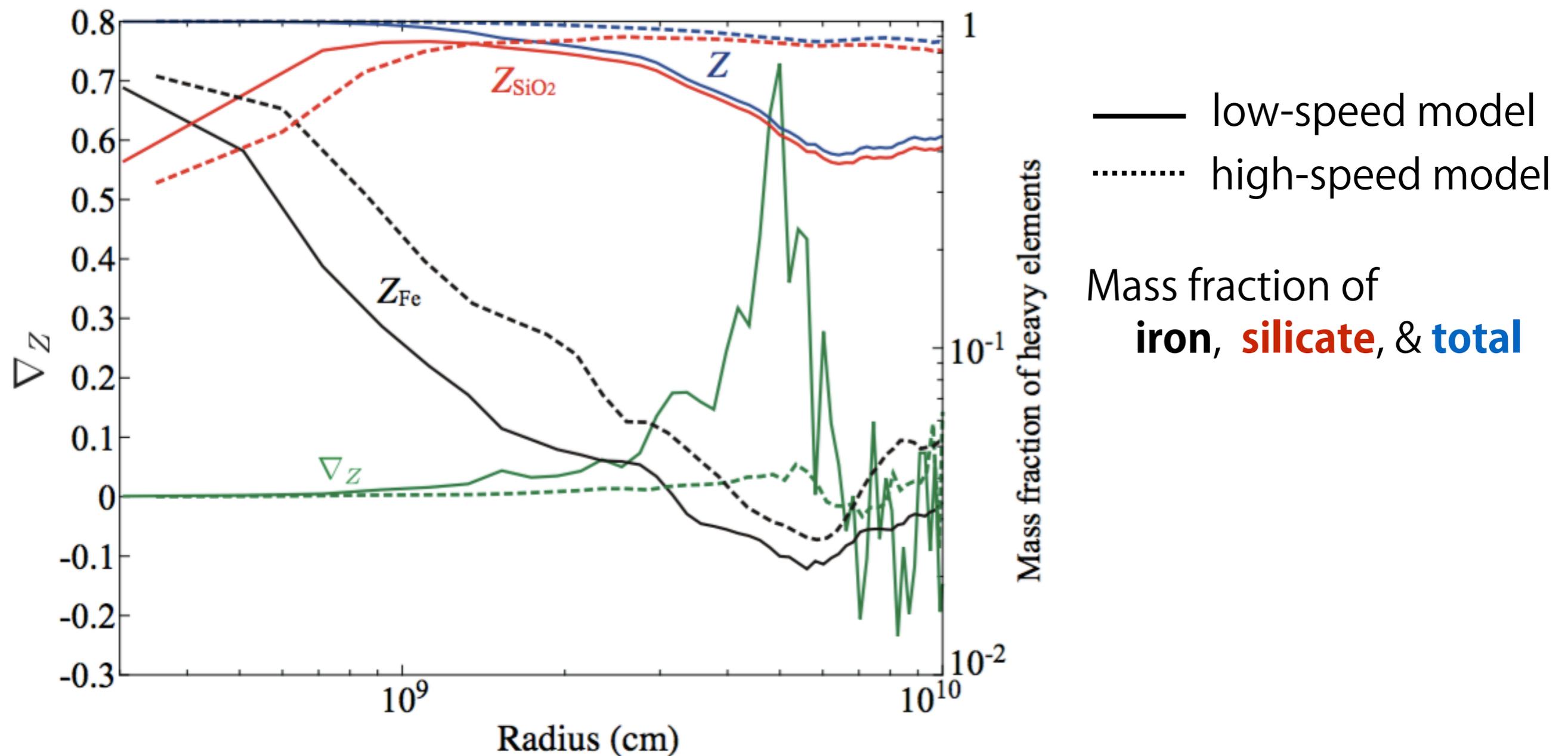


High-speed impact



- An initial layered structure is partly maintained after the collision
- An iron core of the target survives from the impact in both cases and grows in a coalescence manner
- A fraction of rocky material is dredged up in a H/He atmosphere
→ the remaining atmosphere is polluted with heavy elements

Compositional Gradient Inside a Target After an Impact

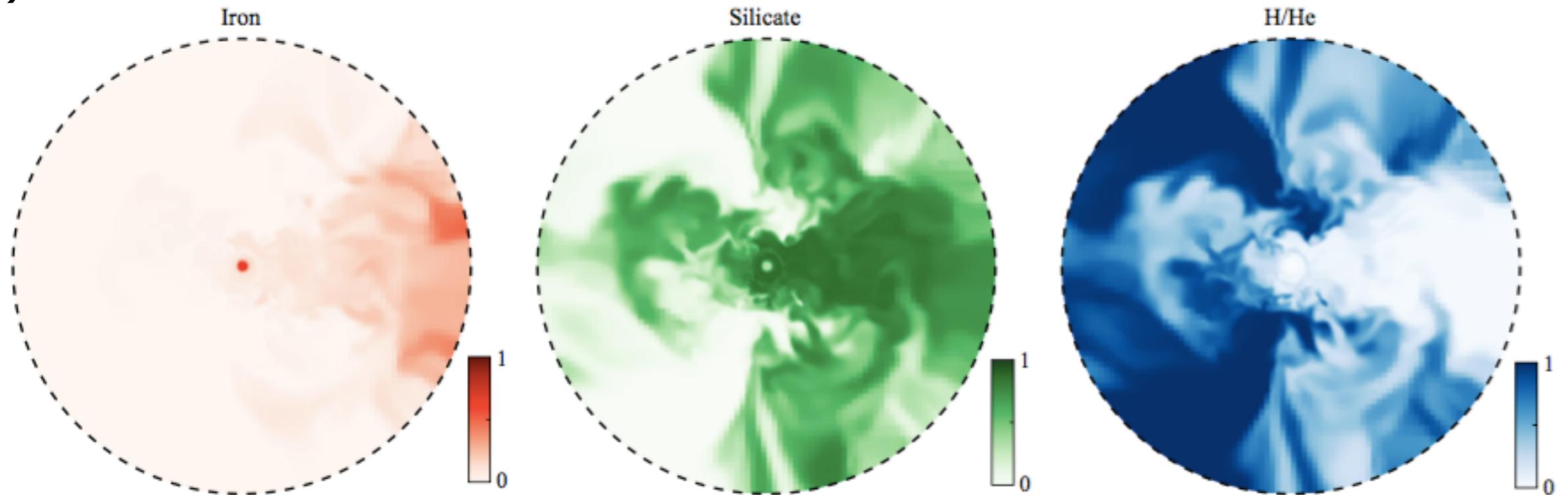


A low-speed head-on collision develops a hot and inhomogeneous interior
→ **a steep, positive compositional gradient** suppresses efficient heat transfer(?)

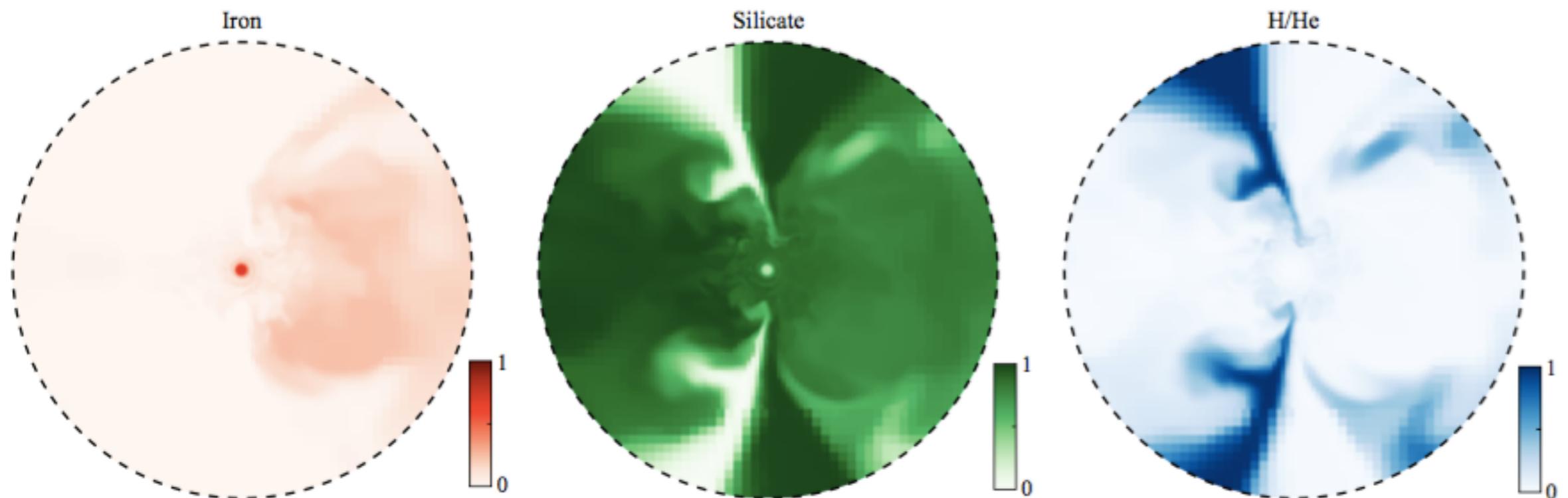
For a high-speed head-on collision,
refractory material is homogenized in the target's interior

Snapshots of Material Mixing After Giant Impacts

(a) 18 hrs after a low-speed impact

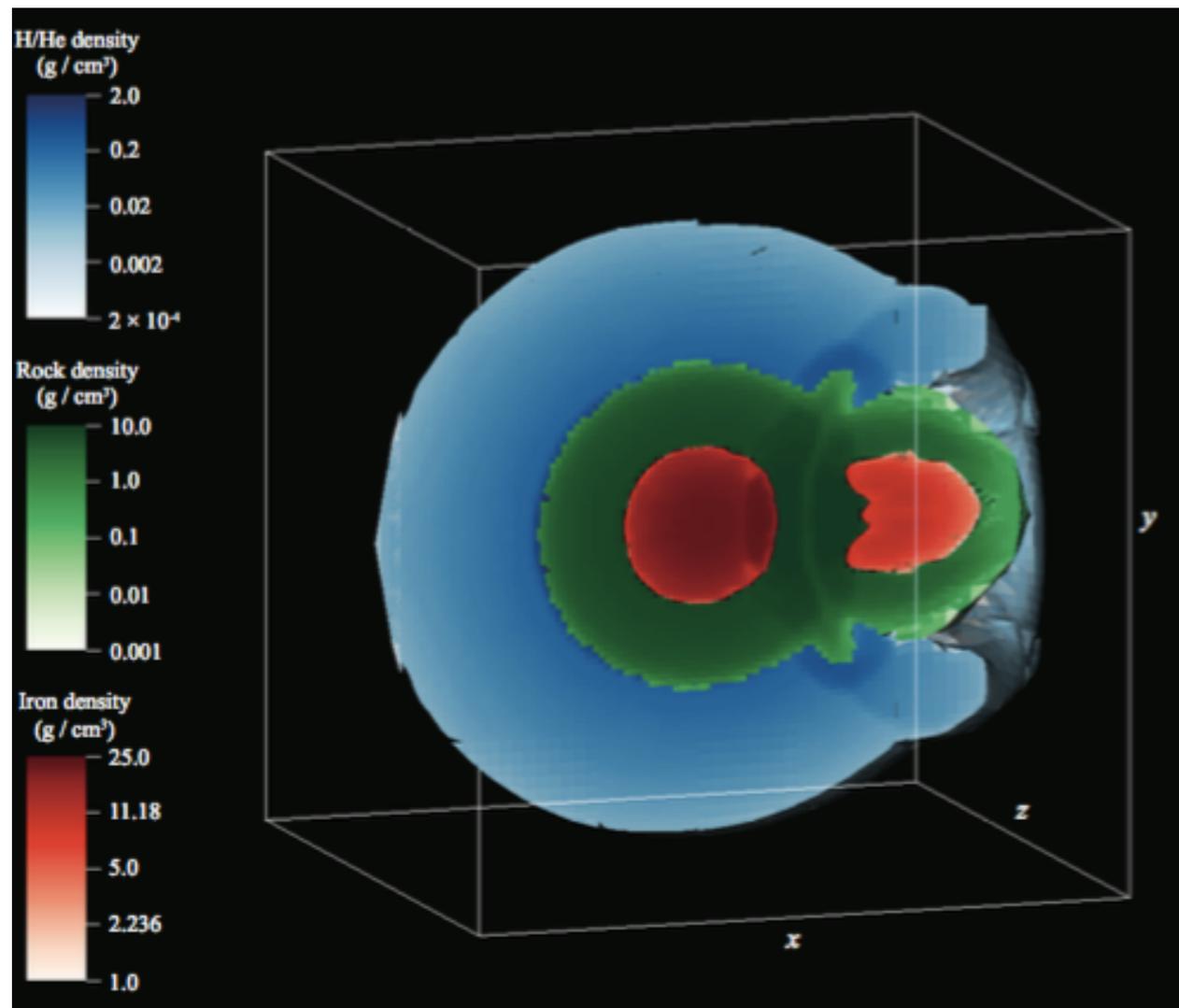


(b) 21.5 hrs after a high-speed impact

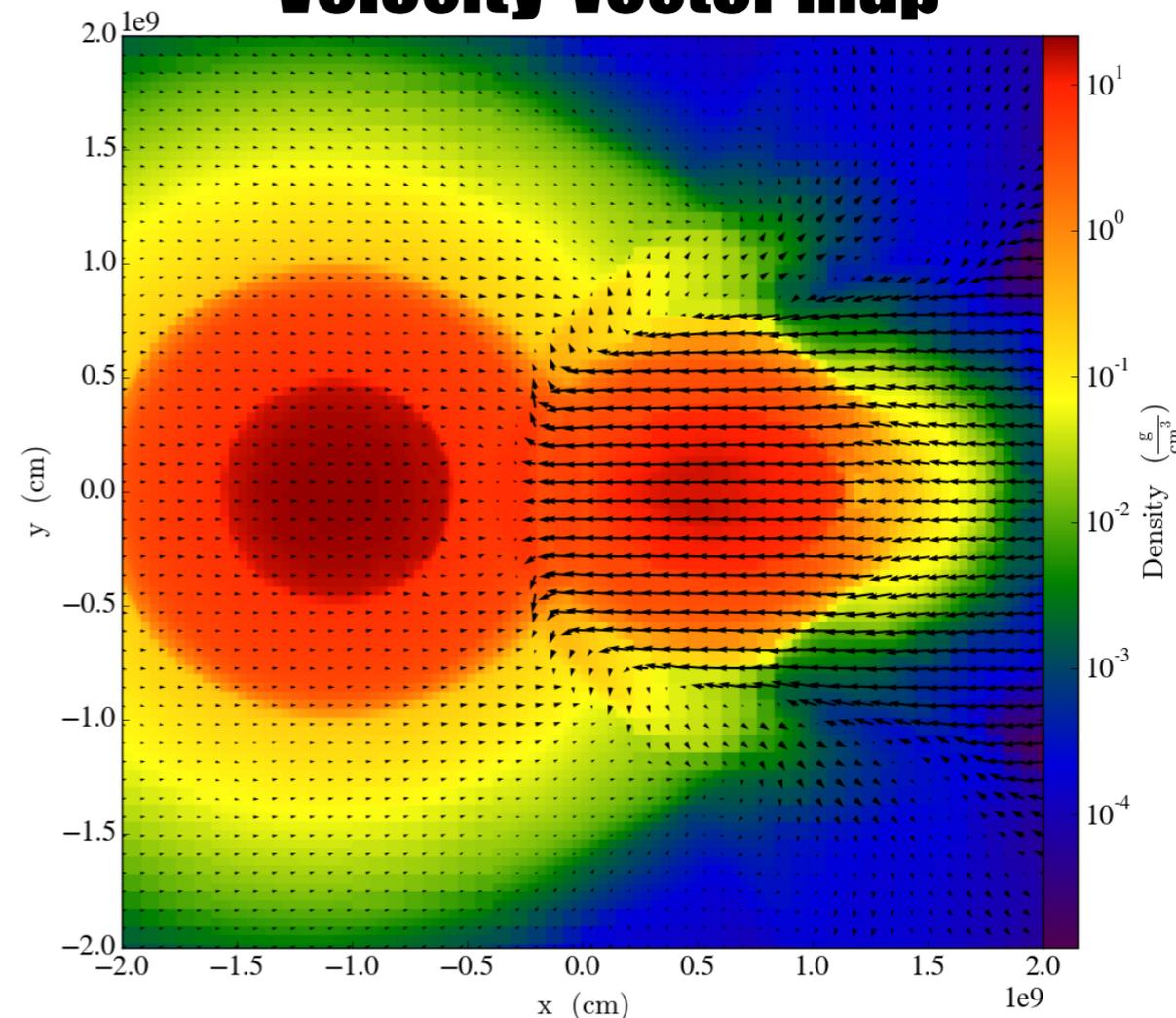


Turbulence or Hydrodynamic Instability?

Species contour



Velocity-vector map



- **A velocity shear** at the interface between two species after an impact → **K-H instability** (at least for short wavelengths)
- **An impact-induced shock wave propagation** → **R-T instability**

However,

An impact-driven turbulence is responsible for **the global mixing**

Take-Home Messages

Different histories of giant impacts result in

- (1) **compositional diversity** of super-Earths (Inadmar & Schlichting, 2015)
- (2) **homogeneous or inhomogeneous** interior
→ suppresses efficient heat transfer
(e.g.) double diffusive convection
- (3) **a hot and inflated atmosphere** (extended beyond the Hill radius)
which **enhances mass loss** via photo-evaporation or a Parker wind
- (4) **the survival of a planetary iron core** through a merger
- (5) **dredge-up of rocky material into a H/He atmosphere** caused by
turbulence driven by an impact-induced shock wave
- (6) **a partial disruption of a three-layered structure**